





# Sensitivity study of the regional climate model RegCM4 to different convective schemes over West Africa. KONE, Brahima<sup>\*1,2</sup>, DIEDHIOU, Arona<sup>1,2,3</sup>, TOURE, N'datchoh Evelyne<sup>1,2</sup>, Sylla, Mouhamadou Bamba<sup>4</sup>, GIORGI, Fillipo<sup>5</sup>, ANQUETIN, Sandrine<sup>2,3</sup>, BAMBA, Adama<sup>1,2</sup>, DIAWARA, Adama<sup>1,2</sup>, and KOBEA, T. Arsene<sup>1,2</sup>

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### BACKGROUNDS

The latest version of RegCM4 with CLM4.5 as a land surface scheme was used to assess the performance and sensitivity of the simulated West African climate system to different convection schemes. The sensitivity studies were performed over the West African domain from November 2002 to December 2004 at a spatial resolution of 50 km × 50 km and involved five convective schemes: (i) Emanuel; (ii) Grell; (iii) Emanuel over land and Grell over ocean (Mix1); (iv) Grell over land and Emanuel over ocean (Mix2); and (v) Tiedtke. It is worth noting that the few previous similar sensitivity studies conducted in the region were performed using BATS as a land surface scheme and involved less convective schemes. Compared with the previous version of RegCM, RegCM4-CLM also shows a general cold bias over West Africa whatever the convective scheme used. This cold bias is more reduced when using the Emanuel convective scheme. In terms of precipitation, the dominant feature in model simulations is a dry bias that is better reduced when using the Emanuel convective scheme. Considering the good performance with respect to a quantitative evaluation of the temperature and precipitation simulations over the entire West African domain and its subregions, the Emanuel convective scheme is recommended for the study of the West African climate system.

## DATA AND METHODOLOGY

Five experiments using the convection schemes of (1) Emanuel over land and Grell over ocean (mix1), (2) Emanuel, (3) Grell, (4) Tiedtke, and (5) Grell over land and Emanuel over ocean (mix2) are conducted using RegCM4-CLM4.5 with 18 sigma levels at 50 km horizontal resolution for the period from November 2002 to September 2004. The first 2 months (i.e., November and December 2002) was considered as spin-up time and not included in the analysis. The years 2003 and 2004 have been selected in this study because they corresponded to a dry and wet year in this region, respectively. The analyses will focus on the rainy season from June to September (JJAS). As quantitative measurements of model skills, we consider mean bias (MB), which is the difference between the area-averaged value of the simulation and the observation, the spatial root mean square difference (RMSD), the spatial correlation called pattern correlation coefficient (PCC) and the distribution of the probability density function (PDF) of the temperature bias. The RMSD, PCC and the PDF provide information at the grid-point level, while the MB does so at the regional level. A Taylor diagram (Taylor, 2001) is used to summarize the assessments above and to show the deviation of different model configuration results from observations. These metrics are computed for each of the subregions indicated in Fig. 1. For this sensitivity study, the model was run at its standard  $con \Box$  guration with 18 vertical sigma layers (model top at 50 hPa) and with initial and boundary conditions provided by the European Centre for Medium-Range Weather Forecasts reanalysis ERA-Interim (Simmons et al., 2007) at a horizontal resolution of 50 km and a temporal resolution of 6 h (00:00, 06:00, 12:00 and 18:00 UTC). Sea surface temperatures (SSTs) were from NOAA optimal interpolation weekly SST data (Reynolds et al., 1994). The terrain characteristics (topography and land use data) were derived from United States Geological Survey (USGS) Global Land Cover Characterization (GLCC; Loveland et al., 2000) at 10 min horizontal resolution. We focus our analysis on the precipitation and the air temperature at 2 m in the summer of June-July-August-September (JJAS) over mainland West Africa. To reduce uncertainty due to lack of surface climate observations over the region (Nikulin et al., 2012), the simulated precipitation is validated using two observational datasets: the GPCP product  $(1^{\circ} \times 1^{\circ})$  resolution) is a satellite derived dataset developed under the Global Precipitation Climatology Project and made available from late 1996 to present and the 0.25° high-resolution dataset of the Tropical Rainfall Measuring Mission 3B43V7 (TRMM) available from 1998 to 2013 (Huffman et al., 2007). The simulated 2 m temperature is also validated using two observational datasets: the Climate Research Unit (CRU) time series version 3.20 gridded at 0.5° of horizontal resolution from the University of East Anglia, available from 1901 to 2011 (Harris et al., 2013), and the University of Delaware version 3.01 (UDEL) gridded dataset at 0.5° of horizontal resolution, available from 1900 to 2010 (Legates and Willmott, 1990). The simulated atmospheric fields are compared with ERA-Interim reanalysis available from 1979 to present at 1.5° of horizontal resolution (Dee et al., 2011). All products have been regridded to  $0.44^{\circ} \times 0.44^{\circ}$  using a bilinear interpolation method to facilitate the comparison with RegCM4 simulations (Nikulin et al., 2012). The model performance is further examined in four subregions (Fig. 1), each with different characteristics of the annual cycle of rainfall: central Sahel (10° W-10°E; 10-16°N), west Sahel (18-10°W; 10-16 °N), Guinea coast (15°W-10 E; 3-10 °N) and West Africa (20° W-20 °E; 5°S-21°N).







